

# A simple model for multiple equilibria in ice-covered oceans

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## Problem formulation

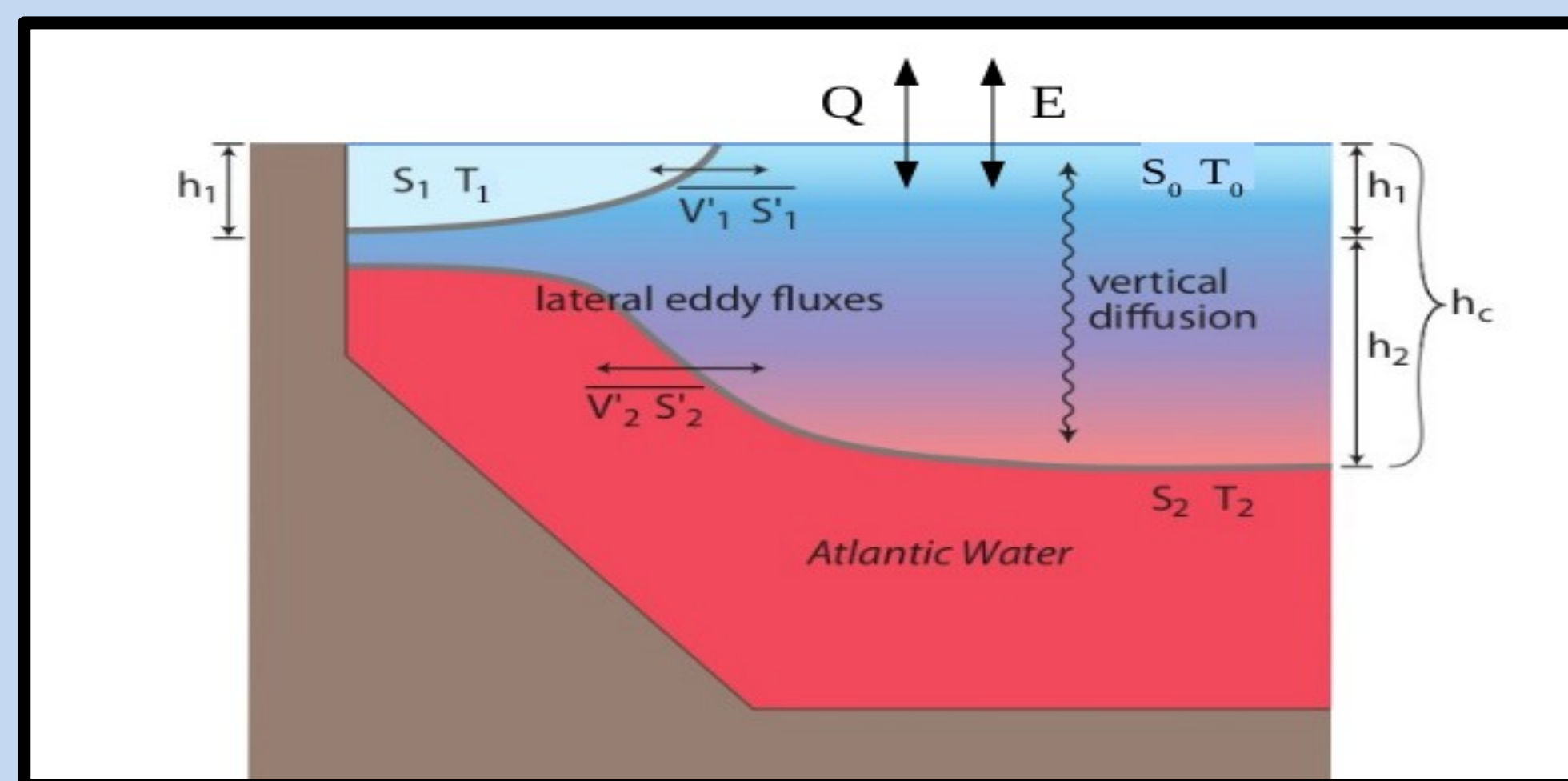
The Arctic Ocean is dominated by a cold, fresh halocline in the upper 50-200 m and a warm, salty water mass of Atlantic origin below. The halocline shields the overlying sea ice from being melted by the warm Atlantic Water

The surface albedo is small for open water and large for ice/snow cover, allowing for possible feedbacks between the surface heat budget and ice cover. (ice-albedo feedback)

Diapycnal mixing in the Arctic is small, in part due to the presence of sea ice. It is thought that mixing would be larger if the Arctic were ice free. (wind-ice-ocean feedback)

We seek guidance on what determines whether multiple equilibria can exist

### A conceptual model for the Arctic Ocean



The salinity and temperature in the basin interior are determined by a balance between lateral eddy fluxes, diapycnal mixing, and air-sea exchange

$$\text{Layers 1+2} \quad Ph_2 \overline{V_2' S_2'} + Ph_1 \overline{V_1' S_1'} = -AES_R$$

$$\text{Layer 2} \quad Ph_2 \overline{V_2' S_2'} = \frac{KA(S_2 - S_0)}{h_1 + h_2}$$

P=perimeter, A=area, K=diapycnal mixing  
E=evaporation - precipitation

The eddy flux is parameterized as

$$\overline{V_k' S_k'} = c_k V_k (S_k - S_0) \quad V_k = \frac{g \alpha_s (S_k - S_0) h_k}{\rho_0 f_0 L}$$

Nondimensionalize variables:

$$\delta S_1 = \frac{S_0 - S_1}{S_2 - S_1} \quad \delta S_2 = \frac{S_2 - S_0}{S_2 - S_1}$$

$$\delta T_1 = \frac{\alpha_T (T_0 - T_1)}{\alpha_S (T_2 - T_1)} \quad \delta T_2 = \frac{\alpha_T (T_2 - T_0)}{\alpha_S (T_2 - T_1)}$$

## Analytic solutions

We can derive equations for the interior salinity  $\delta S_1$  and  $h_2$

$$(\delta S_1^2 - \phi)^{3/2} - \lambda (1 - \delta S_1)^2 = 0 \quad h_2 = (c_1/c_2)^{0.5} \frac{(\delta S_1^2 - \phi)^{1/2}}{1 - \delta S_1}$$

Solve for temperature subject to a surface heat flux

$$Q = G(T_0 - T_A) + (1 - \alpha) Q_S$$

$T_A$  is atmospheric temperature

G is restoring constant ( $W/m^2 C$ )

$\alpha$  = albedo (0.6 ice cover, 0.1 open ocean)

$Q_S$  is solar radiation

We can also derive a solution for the temperature of the halocline

$$\delta T_2 = \frac{\gamma \delta T_A + \delta S_1 \delta T_B - (1 - \alpha) Q_S}{\gamma + \delta S_1 + c_1/c_2 h_2^2 (1 - \delta S_1)}$$

### Nondimensional parameters

$$\phi = \frac{AES_R}{c_1 \psi_g (S_2 - S_1)} \quad \gamma = \frac{AG}{c_1 \rho_0 C_p \psi_g} \quad \lambda = \frac{c_2^{0.5} KA}{c_1^{1.5} \psi_g h_1}$$

evap minus precip                      cooling                      mixing

$$\psi_g = \frac{gP \alpha_s (S_2 - S_1)}{\rho_0 f_0 L} \quad Q_S = \frac{\alpha_T A Q}{c_1 \rho_0 C_p \alpha_s (S_2 - S_1)}$$

reference transport                      solar radiation

$$\delta T_B = \frac{\alpha_T (T_2 - T_1)}{\alpha_S (S_2 - S_1)} \quad \delta T_A = \frac{\alpha_T (T_2 - T_A)}{\alpha_S (S_2 - S_1)}$$

boundary temperature                      atmospheric temperature

### For typical parameters:

$$\lambda \ll 1$$

weak mixing

$$\gamma = O(1)$$

cooling

$$Q_S = O(10^{-1})$$

solar radiation

$$\lambda \gg 1$$

strong mixing

$$\phi = O(10^{-2})$$

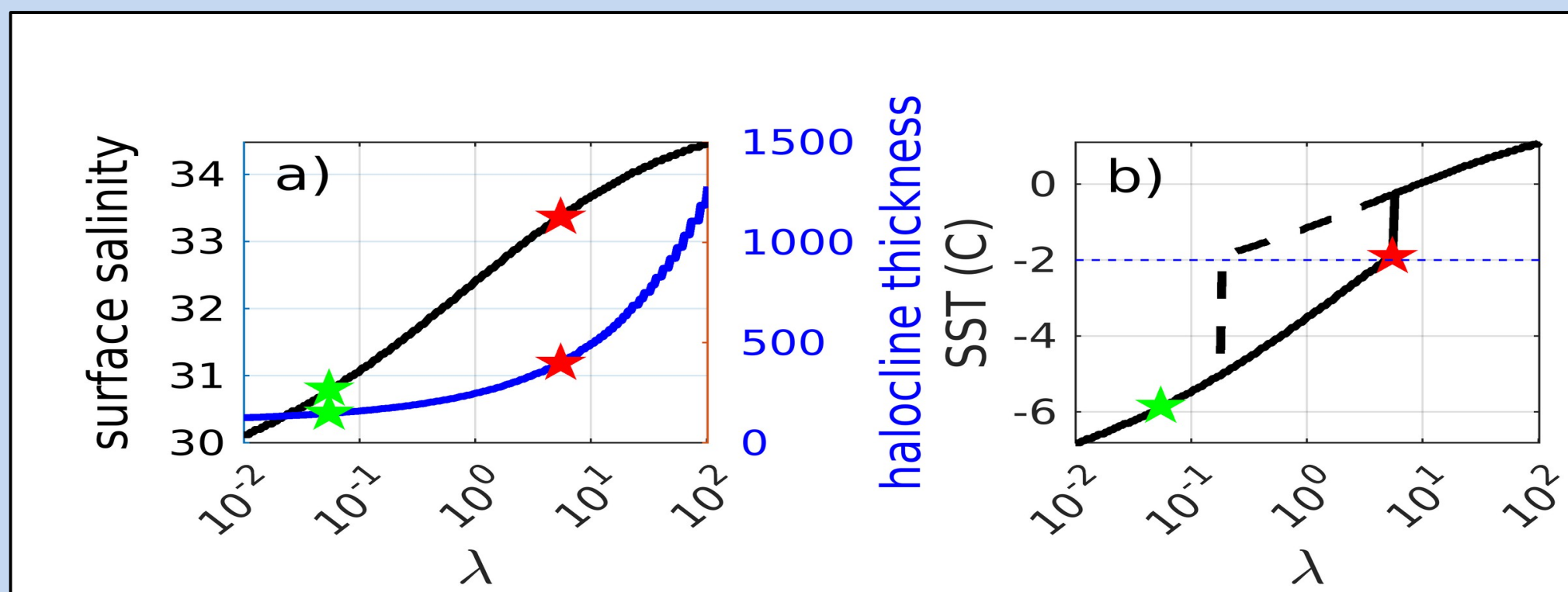
precipitation

$$\delta T_A = O(10^{-1})$$

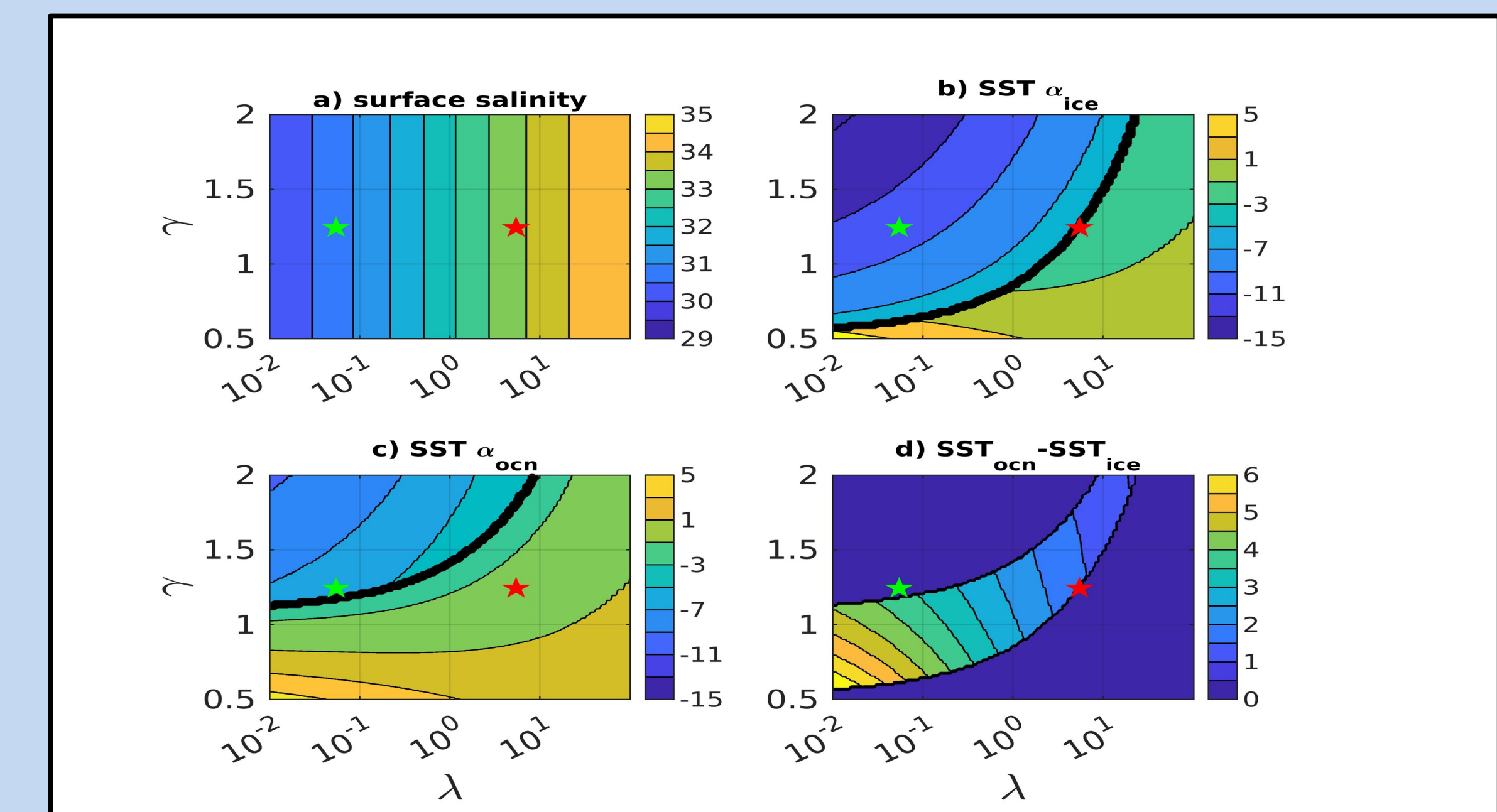
atmospheric temp

## Results

### Dependence of halocline on mixing



Salinity, halocline depth, and temperature increase with increasing mixing. Multiple equilibria are possible, one state with ice-covered with high albedo and one ice-free with low albedo (ice-albedo feedback)



ME exist over a finite range of cooling, most prevalent for weak mixing  
Also get ME if mixing gets large when the ice melts (wind-ice-ocean feedback)

Similar analysis can be done for all forcing parameters...

### SUMMARY

Coupled equations for the temperature, salinity, and halocline depth have been derived that describe a balance between lateral eddy fluxes, vertical mixing, and air-sea exchange

The system supports multiple equilibria resulting from:

- > The dependence of albedo on ice cover
- > The expected increase in mixing as the Arctic becomes ice-free
- > Asymptotic solutions give explicit expressions for weak and strong mixing regimes (ice-covered and ice-free), including the level of mixing required to melt the ice
- > Existence of ice cover depends on many parameters
  - External forcing (atmospheric temperature, E-P, cooling rate, solar, source waters)
  - Environmental parameters of the basin (e.g. size, shelf depth, shelf width, mixing rate)